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National contributions to climate change due to historical emissions of carbon dioxide, methane, and nitrous oxide since 1850

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Anthropogenic emissions of carbon dioxide (CO₂), methane (CH₄) and nitrous oxide (N₂O) have made significant contributions to global warming since the pre-industrial period and are therefore targeted in international climate policy. There is substantial interest in tracking and apportioning national contributions to climate change and informing equitable commitments to decarbonisation. Here, we introduce a new dataset of national contributions to global warming caused by historical emissions of carbon dioxide, methane, and nitrous oxide during the years 1851–2021, which are consistent with the latest findings of the IPCC. We calculate the global mean surface temperature response to historical emissions of the three gases, including recent refinements which account for the short atmospheric lifetime of CH₄. We report national contributions to global warming resulting from emissions of each gas, including a disaggregation to fossil and land use sectors. This dataset will be updated annually as national emissions datasets are updated.

Background and Summary

Anthropogenic emissions of carbon dioxide (CO₂), methane (CH₄) and nitrous oxide (N₂O) are key components responsible for climate change since the pre-industrial period^{1,2}. The use of fossil carbon sources in the energy, industry, transport, waste and product use sectors and land use, land use change and forestry (LULUCF), has led to increased atmospheric concentrations of CO₂, CH₄ and N₂O and driven Earth's surface energy balance into surplus^{3–6}. The Intergovernmental Panel on Climate Change (IPCC) estimated in its Sixth Assessment Report (AR6) that rising atmospheric concentrations of CO₂, CH₄ and N₂O have already caused global mean surface temperature (GMST) to increase by 1.4 °C (0.9–2.2 °C at the 90% confidence interval) in the industrial era, independent of other greenhouse gases (GHGs), ozone precursors (e.g., VOC, CO, NO_x), and aerosols (e.g., SO₂, black carbon and organic carbon) which either heat or cool the planet for a net warming of 1.1 °C attributable to human activities^{1,2,7,8}.

Due to their long-lived or powerful effects on climate, the national emissions of CO₂, CH₄ and N₂O are widely regulated by the United Nations Framework Convention on Climate Change (UNFCCC). All parties to the convention set targets for CO₂ in the form of nationally determined contributions (NDCs)⁹, while around 90% of NDCs include targets for CH₄ and N₂O. Consequently, keeping track of the emissions of CO₂, CH₄ and N₂O and the climatic responses to those emissions is especially important for ensuring accountability with respect to NDCs. The current work seeks to inform the 2023 Global Stocktake of the UNFCCC, the formal process by which national progress on NDCs is evaluated. While we focus on three CO₂, CH₄ and N₂O that are

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included in most NDCs, we note that future work should seek to include other important GHGs, such as fluorinated gases (F-gases), which are also included in the NDCs of some countries.

The contributions of individual countries to the warming caused by rising CO₂, CH₄ and N₂O concentrations has changed through time and depends strongly on the unique history of land clearing, industrialisation, and decarbonisation of a country. Emissions of fossil CO₂ – the largest driver to historical climate change¹ – have generally continued to rise with economic growth in developing countries even after the establishment of the UNFCCC in 1992. On the other hand, fossil CO₂ emissions in developed countries have begun to decline after increasing throughout much of the industrial era since the mid-19th Century^{3,10,11}. National contributions to climate change are closely tied to cumulative emissions of CO₂ in the industrial era because a substantial fraction of emitted CO₂ remains in the Earth's atmosphere for centuries^{6,12}. Consequently, emissions from developed nations have contributed significantly to warming since the industrial revolution^{12,13}. Tracking national contributions to climate change is thus critical to understanding the burden of responsibility that a country carries for global warming and can further inform the design of international policies that pursue equitable decarbonisation pathways.

Numerous studies have assessed national contributions to the historical change in GMST caused by anthropogenic emissions of CO₂ and other gases or aerosols. These studies have typically employed simple climate models to simulate the GMST response to the emissions of CO₂^{14–17}, CH₄ and N₂O^{18–20}, and other gases and aerosols^{12,13,21–26}. Several studies have further evaluated how national contributions to warming may develop in future periods based on current pledges, nationally declared contributions or emission scenarios^{17,25,27}. Others have employed coupled climate models to evaluate national contributions to simulated climate change impacts resulting from warming, such as the likelihood of extreme weather, sea level rise, and economic growth^{17,20,28,29}. Within these studies, each climate model or ensemble of models exhibits a specific climate sensitivity such that the response of GMST to anthropogenic emissions is model- or ensemble-specific³⁰.

Recent advances in the understanding of the response of GMST to cumulative CO₂ emissions have provided express formulas that can be used to efficiently calculate the warming resulting from cumulative emissions of CO₂^{6,31–34}. The transient climate response to cumulative emissions of CO₂ (TCRE), a coefficient representing the warming per trillion tonnes of carbon emitted as CO₂, was estimated in the IPCC's Sixth Assessment Report to be 0.45 °C per 10³ Pg CO₂ (0.3–0.6 °C per 10³ Pg C at the 90% confidence interval) based on synthesis of estimates made in 28 earlier studies (see **Methods**; Eq. (1)). In addition, global warming potential (GWP) metrics have been developed to relate emissions of non-CO₂ greenhouse gas emissions to cumulative CO₂ emissions, most recently including the GWP* approach which captures the contrasting impacts of short and long-lived climate pollutants on GMST^{35–39}. An advantage of using express formulas based on TCRE and GWP* coefficients to calculate the warming response to emissions is that the values of each coefficient can be selected in line with IPCC best-estimates and uncertainty ranges, which synthesise the behaviour shown by many different models across dozens of studies^{3,6}. Hence, the climatic response to emissions of GHGs does not depend on the climate sensitivity of an individual model. Other straightforward equations (see **Methods**; Eqs. (2, 3)) have been devised to relate the warming effect of CH₄ and N₂O emissions to the warming effect of cumulative CO₂ emissions over defined time horizons, often 100 years^{36–38}. The express equations are yet to be applied to estimate the combined effects of the national emissions of multiple gases on GMST.

Here, we present a new dataset of changes in GMST during 1851–2021 resulting from historical emissions of CO₂, CH₄ and N₂O at the global scale and for individual countries. Our estimates of warming are based on the application of express equations for TCRE and Global Warming Potential (GWP) to the emissions time series for each gas, using current best-estimates of the coefficients involved in the equations from IPCC AR6 (see **Methods**)^{3,40–43}. We further estimate the contributions of fossil and LULUCF emissions to change in GMST at global and national levels. To enhance the relevance of the dataset to international organisations, we also provide estimates of the contributions of various country groupings including Annex I countries (number of countries, n = 42), Annex II countries (n = 23), economies in transition (EITs; n = 15), the least developed countries (LDCs; n = 47), and like-minded developing countries (LMDC; n = 24) as defined by the UNFCCC. We also consider the contributions of the organisation for economic co-operation and development (OECD; n = 38), the European Union (EU27 post-Brexit), and the Brazil, South Africa, India and China (BRICS) group. Lists of the countries included in each country grouping are provided with the **Data Records**⁴⁴.

This dataset is publicly available via a Zenodo repository⁴⁴. For completeness, we provide the annual and cumulative emissions data used to calculate the change in GMST. We emphasise that the emissions data are not introduced here but rather derive from the Global Carbon Project (GCP)^{3,40} and the national historical emissions time series component of the Potsdam Realtime Integrated Model for probabilistic Assessment of emissions Paths dataset (PRIMAP-hist)^{41,42}. We focus on the period since 1850 to align with the period over which warming was usually evaluated in the IPCC AR6⁴⁵. The final year of 2021 is determined by the mutual period covered by the requisite datasets. We will update our dataset annually in line with updates and extensions to the GCP^{3,40} and PRIMAP-hist^{41,42} emissions datasets.

Methods

Emissions time series. *Carbon dioxide.* We retrieved estimates of territorial fossil CO₂ emissions for the years 1850–2021 from the GCP, as published in its 2021 assessment of the global carbon budget (GCB)^{3,40}. The GCB estimates include national annual emissions of CO₂ due to coal, oil and natural gas combustion, the use of these fuels in non-combustive industrial processes, and the production of cement clinkers^{3,40}. The emissions relate to the energy, industry, transport, product use, solvent use and waste sectors. Under the GCB methodology, national CO₂ emissions are preferentially taken from the country submissions to the United Nations Framework Convention on Climate Change (UNFCCC) for 42 Annex I countries. Emissions in other countries and in Annex I countries prior to 1990 predominantly derive from the Carbon Dioxide Information Analysis

Center (CDIAC)⁴⁶. For the years since 2020, not included in either the UNFCCC or CDIAC datasets, the national emissions are estimated using national or regional energy growth rates from the annual BP Statistical Review of World Energy⁴⁷. Cement emissions are based on national inventories of cement production and ratios of clinker production from officially reported clinker production data and emission factors, IPCC default emission factors, industry-reported clinker production, and survey-based clinker ratios⁴⁸. A more complete description of methodology used to compile the fossil CO₂ emissions time series is provided in our previous work (refs. 3,40).

In addition, we retrieved estimates of historical CO₂ emissions from LULUCF for the years 1850–2021 from the average of the three bookkeeping models that contribute to the GCB (BLUE, Houghton & Nassikas [H&N], and OSCAR)^{49–51}. The bookkeeping models combine historical estimates of changes in agricultural and forest areas, other loss of primary vegetation and wood harvesting with response curves that describe the decay of vegetation and soil carbon over time, including transfer to product pools of different lifetimes, as well as carbon uptake due to regrowth. Emissions from peat burning and drainage are added from external datasets^{52–55}. This way gross emissions and removals due to deforestation, afforestation, logging and forest degradation (including harvest activity), shifting cultivation (cycle of cutting forest for agriculture, then abandoning), conversion of pastures and grasslands, and regrowth of forests following wood harvest or abandonment of agriculture are captured³. A key difference between the bookkeeping estimates is their source of historical land-use data: gridded data from the Land Use Harmonization dataset (LUH2)⁵⁶ in the case of BLUE; national data from the Forest Resource Assessment (FRA)⁵⁷ and agricultural areas⁵⁸ from the United Nations Food and Agriculture Organisation (FAO) in H&N, and; a combination of LUH2 and FRA data in OSCAR). All datasets are ultimately based on information of FAO agricultural area and national wood harvest statistics, both available since 1961, or forest area, available since 1990. The information is extended to earlier time periods using other sources of information such as population growth or regional historical sources^{50,59}. While the minimum spatial units in H&N and OSCAR are countries or country groupings, simulations with BLUE are spatially explicit at quarter-degree resolution and carbon fluxes aggregated afterwards to the country level.

Using the abovementioned datasets, we constructed a time series of cumulative total, fossil and LULUCF CO₂ emissions for the years 1851–2021 (i.e. since the base year of 1850). Estimates are provided for each country and country grouping as well as for the global total.

Methane and nitrous oxide. We retrieved estimates of historical N₂O emissions for the years 1850–2021 and CH₄ emissions for the years 1830–2021 from the Potsdam Realtime Integrated Model for probabilistic Assessment of emissions Paths dataset (PRIMAP-hist version 2.4)^{41,42}. PRIMAP-hist combines various datasets into a global multi-gas emissions record. For CH₄, the data for years 1830–1849 are required to estimate warming for the years 1850 through 1869 (see **Non-CO₂ Short-lived Climate Forcers**, below).

For CH₄ and N₂O, the fossil emissions estimates in PRIMAP-hist relate to emissions from energy, industry, transport, product use, solvent use and waste. LULUCF emissions in PRIMAP-hist HISTTP include emissions caused by land use changes, such as deforestation or agricultural land abandonment, and agricultural emissions of N₂O and CH₄ including various sources, such as rice cultivation, synthetic fertilizers, and manure management (sectors ‘livestock’ and ‘agriculture excluding livestock’). PRIMAP-hist adopts data from various underlying datasets from the UNFCCC⁶⁰, the Food and Agriculture Organization (FAO)^{42,58}, the Emissions Database for Global Atmospheric Research (EDGAR)^{61,62}, and the Community Emissions Data System (CEDS)⁶³. Details of the prioritisation given to the various data sources included in PRIMAP-hist are described in ref. 41. We selected the HISTTP scenario of the PRIMAP-hist dataset, in which emissions estimates based on third-party datasets (e.g. research institutes, international organisations, or private companies) are prioritized over country-reported data, rather than the HISTCR scenario in which country-reported data are prioritised over third-party datasets. This selection is made because the estimates of the PRIMAP-hist HISTTP scenario are closer to the average of a wider range of emissions products than the HISTCR scenario during the period 1970–2018⁶⁴.

Using the abovementioned datasets, we constructed a time series of cumulative total, fossil and LULUCF emissions of CH₄ and N₂O for each year during 1851–2021 (since the base year 1850), all expressed in CO₂-equivalent terms (see section ‘CO₂-equivalent Emissions of Non-CO₂ Species’). Estimates are provided for each country and country grouping as well as for the global total.

Global mean surface temperature response to cumulative CO₂ emissions. In recent years, methods for the express calculation of the GMST responses to cumulative emissions of CO₂ and non-CO₂ gases have emerged. A range of studies demonstrated that the CO₂-induced GMST increase is proportional to cumulative carbon emissions of CO₂, regardless of the path taken to reach peak cumulative CO₂ emissions^{6,31–34}. Consequently, a constant known as the transient climate response to cumulative emissions of CO₂ (TCRE) can be defined as the change in GMST per 1,000 Pg C emitted (°C per 10³ Pg C). In its AR6⁶, the IPCC synthesised estimates of TCRE from dozens of model and observation-based studies of TCRE and arrived at the current best estimate of 1.65 °C per 10³ Pg C emitted (likely range 1.0–2.3 °C per 10³ Pg C). The GMST response to cumulative emissions of CO₂ can be expressly calculated as follows:

$$\Delta T = \kappa \times \frac{1}{C} \times E_{CO_2} \quad (1)$$

where ΔT is the change in GMST (°C), k is the TCRE (°C per 10³ Pg CO₂ emitted), E_{CO_2} is the cumulative emissions of CO₂ during the period of interest (10³ Pg CO₂), and C is a constant used to convert the mass of carbon in a CO₂ molecule to the total mass of a CO₂ molecule ($C = 3.664 \text{ Pg CO}_2 \text{ Pg C}^{-1}$)³. Note that the IPCC’s central estimate for TCRE in CO₂ terms ($\kappa \times \frac{1}{C}$) is 0.45 °C per 10³ Pg CO₂ emitted.

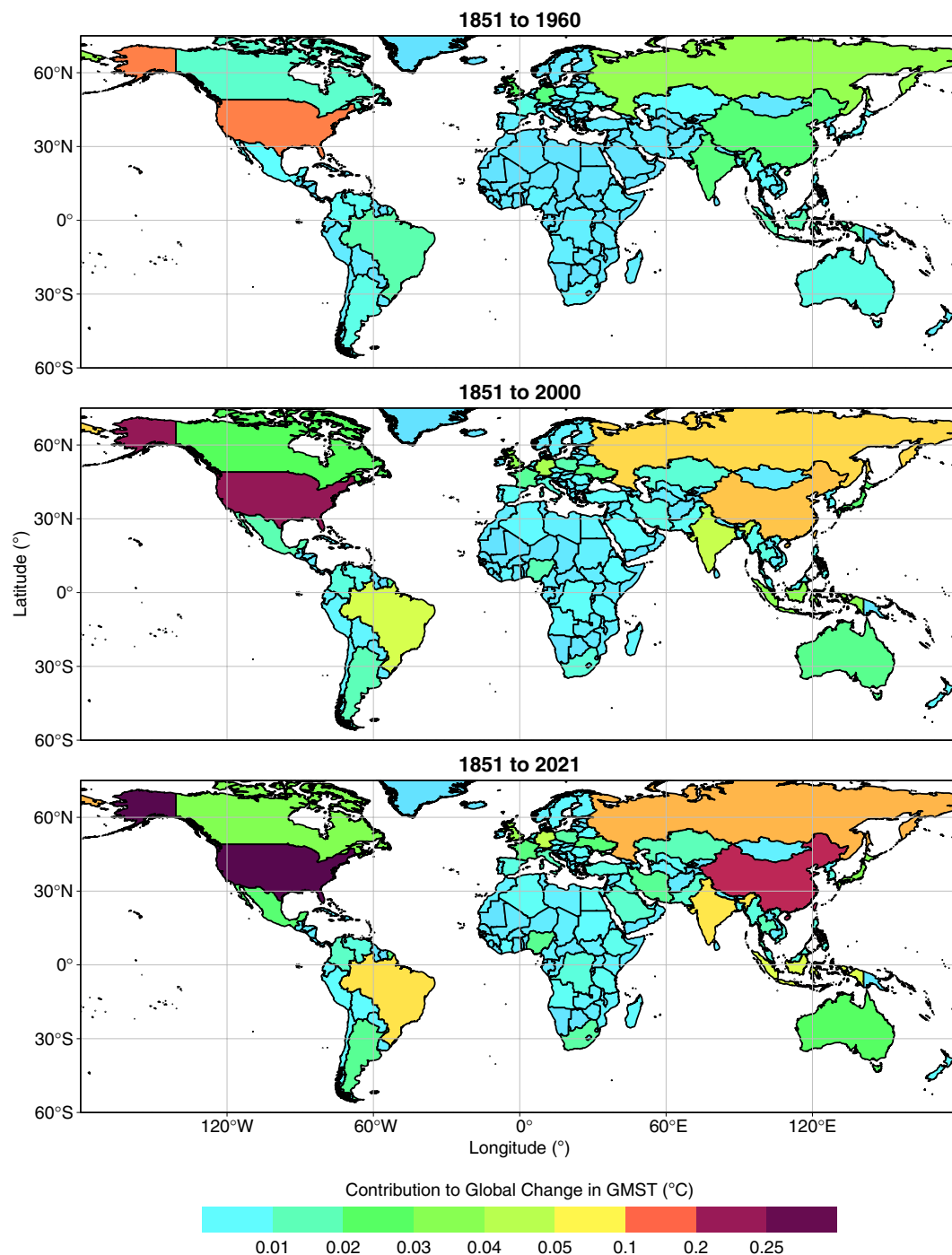


Fig. 1 National contributions to change in global mean surface temperature (GMST, °C) resulting from historical emissions of CO₂, CH₄ and N₂O during three time periods. All data shown are provided in the **Data Records**⁴⁴.

The TCRE has previously been used to make express calculations of the cumulative emissions of CO₂ that remain to be emitted before GMST exceeds a chosen warming target (e.g. 1.5°C or 2.0°C), known as the remaining carbon budget^{65–68}. For example, the remaining carbon budget is regularly re-assessed in the United Nations Environment Programme’s Gap Report⁶⁹. It has also been used in previous work to expressly calculate national contributions to warming caused by CO₂ emission¹².

Here, we use Eq. (1) to estimate the GMST response to cumulative CO₂ emissions in years 1851–2021 (since the base year 1850), adopting the IPCC best-estimate for TCRE (0.45°C per 10³ Pg CO₂).

CO₂-equivalent emissions of Non-CO₂ species. *Non-CO₂ long-lived climate forcers (N₂O).* The GMST response to emissions of other long-lived climate forcers (LLCFs) can be estimated by first expressing their

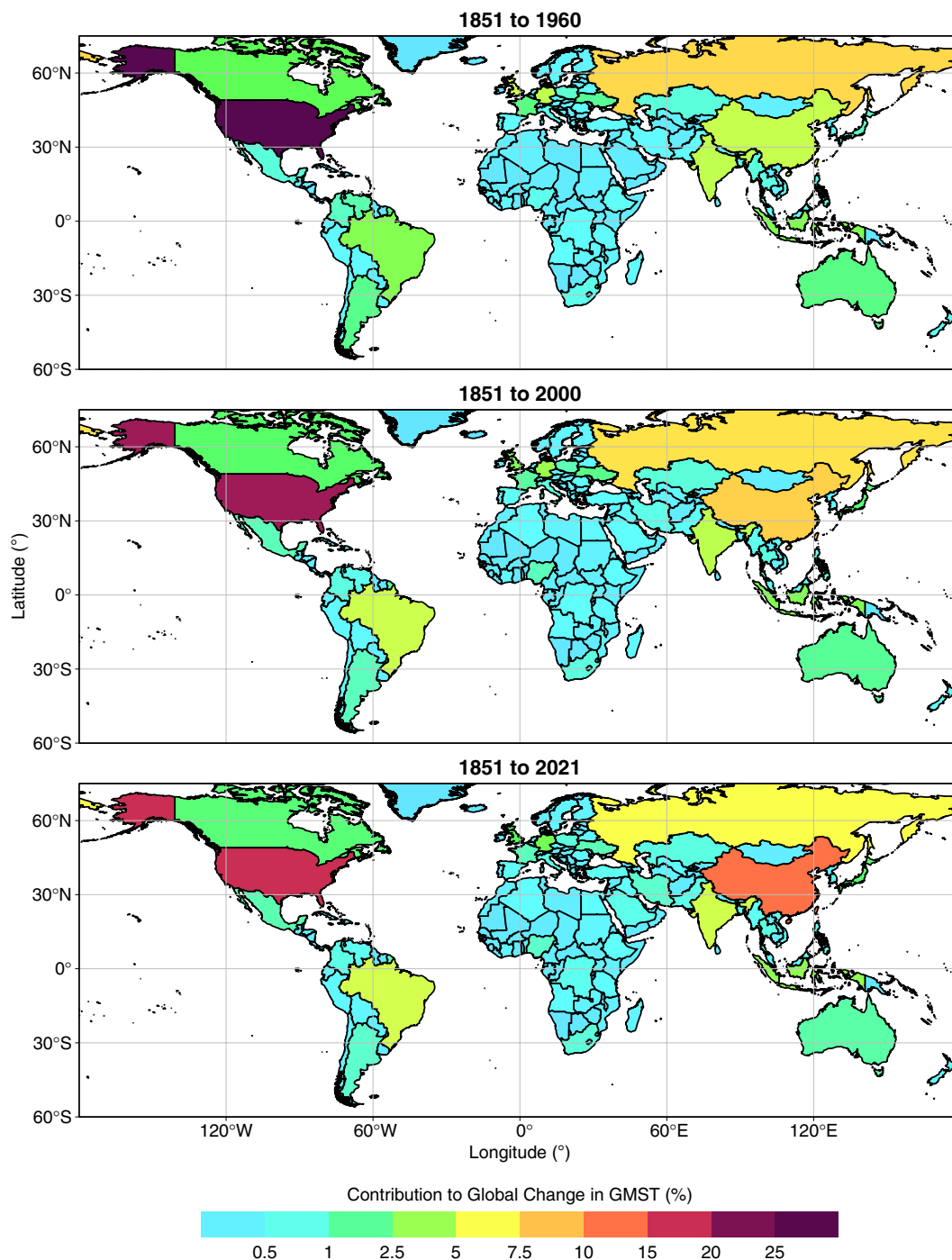


Fig. 2 National contributions to change in global mean surface temperature (GMST) resulting from historical emissions of CO₂, CH₄ and N₂O during three time periods. Values are expressed as a percentage of the change in GMST due to all global emissions of the three gases. All data shown are provided in the **Data Records**⁴⁴.

emissions in terms of the cumulative CO₂ emissions that would result in equivalent warming over a selected time horizon. For conversion of cumulative non-CO₂ LLCF emissions to an equivalent quantity of cumulative CO₂ emissions, one widely used metric is the global warming potential (GWP)². GWP expresses the time-integrated radiative forcing (warming) caused by a pulse emission of a non-CO₂ greenhouse gas, relative to a pulse emission of an equal mass of CO₂ (refs. ^{70,71}). The GWP of a gas depends on its radiative efficiency (infrared energy absorbed) relative to that of CO₂, which has been determined using spectroscopy across a range of gas mixing ratios relevant to the study of Earth's climate⁷⁰. In addition, the GWP of a gas depends on its atmospheric lifetime and the time horizon of interest. As the residence time of a non-CO₂ GHG in the atmosphere can be longer (or shorter) than that of CO₂, the proportion of the non-CO₂ gas that remains in the atmosphere at any time following the pulse emission can be greater (or smaller) than the proportion of CO₂ that remains in the atmosphere and

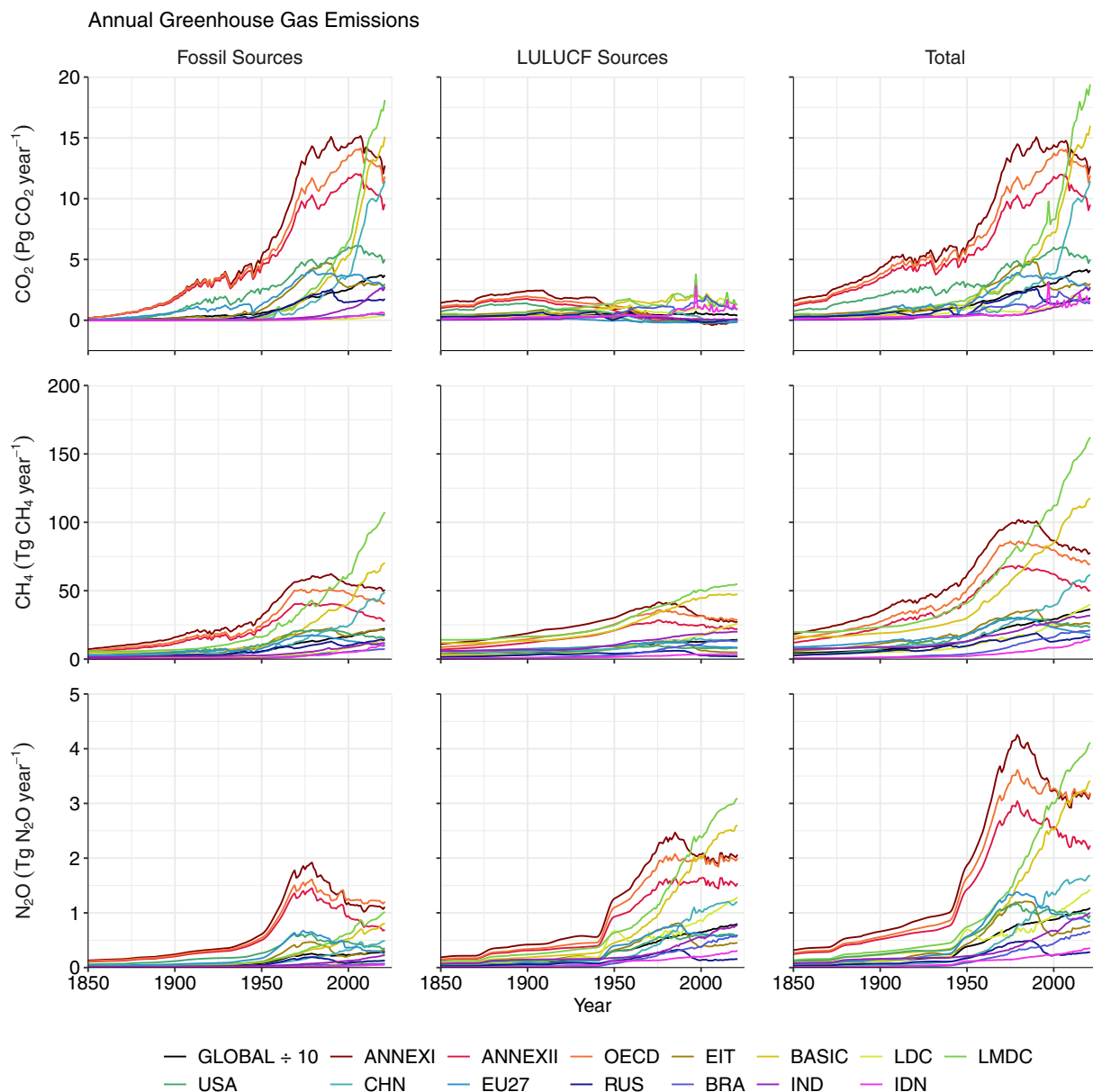


Fig. 3 Annual emissions of CO₂ (Pg CO₂ year⁻¹), CH₄ (Tg CH₄ year⁻¹) and N₂O (Tg N₂O year⁻¹) during 1850–2021, shown globally and for selected countries and country groupings. The primary data sources are the global carbon budget and PRIMAP-hist^{3,40–42} (see **Methods** for data sources). The ISO3 codes of countries in the legend are: CHN, China; RUS, Russia; BRA, Brazil; IND, India; IDN, Indonesia. All data shown are provided in the **Data Records**¹⁴.

the original pulse emission can thus have a more (or less) lasting impact on the energy balance of the atmosphere. Consequently, the GWP metric is dependent on time since emission; its value at a relatively near time horizon (e.g. 20 years) can differ from its value at a later time horizon (e.g. 100 years).

GWP values calculated over a time horizon of 100 years (denoted GWP₁₀₀) have been employed particularly extensively in climate studies to report emissions of various GHGs on the same scale as CO₂ (ref. ²). GWP₁₀₀ values are used to estimate cumulative CO₂-equivalent emissions of an LLCF over a 100-year time horizon, denoted $E_{CO_2-e_{100}}$ (Pg CO₂-e₁₀₀) in Eq. (2):

$$E_{CO_2-e_{100}} = E_{LL} \times GWP_{100} \quad (2)$$

where E_{LL} is the cumulative emissions of the LLCF (Pg LLCF) and GWP₁₀₀ is a constant representing the mass of CO₂ that would result in equivalent warming over a 100-year time horizon (unitless). For example, the current best estimate for the GWP₁₀₀ of N₂O is 273 (ref. ²), signifying that 1 Pg of N₂O results in the same warming over a 100-year horizon as 273 Pg of CO₂.

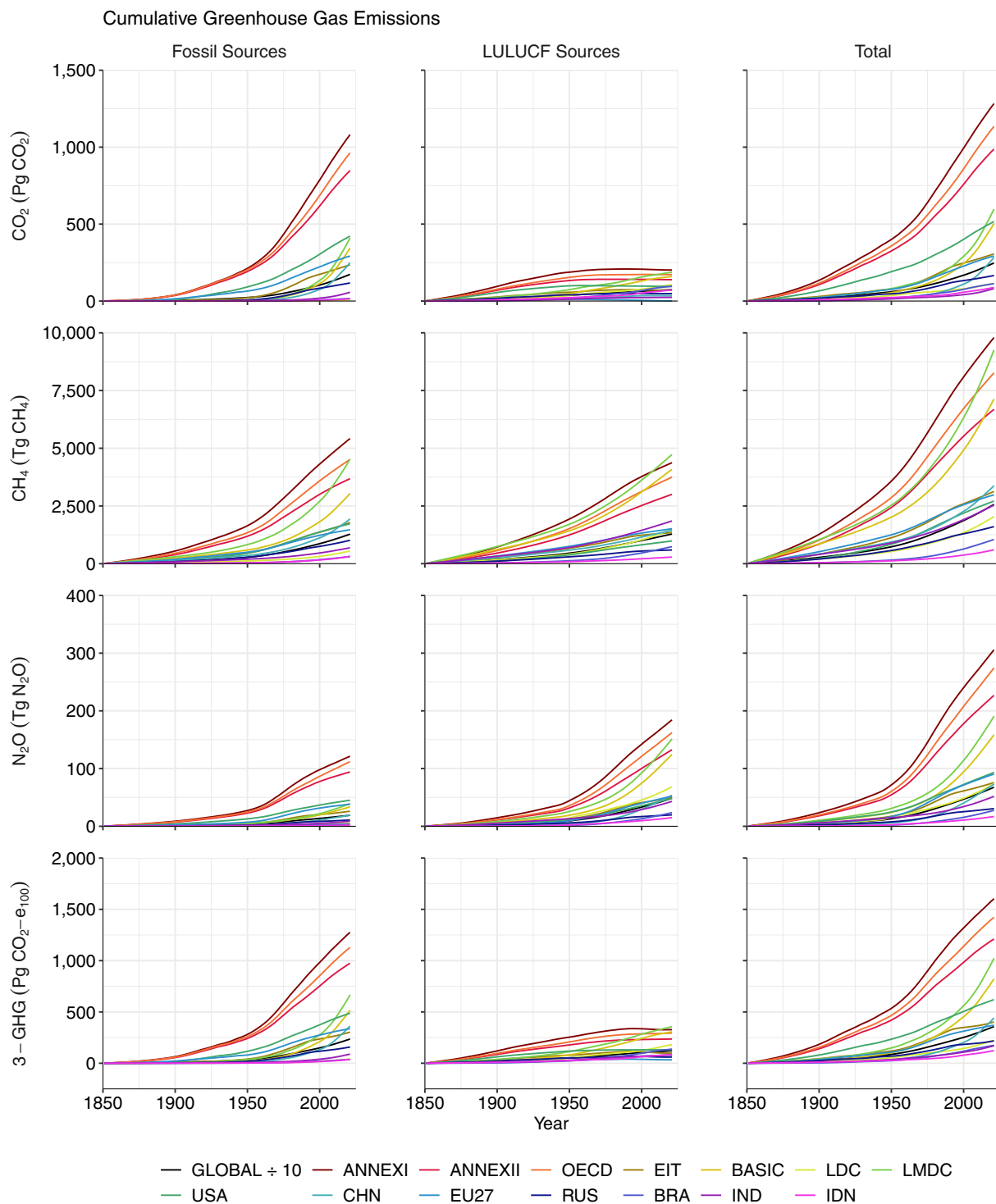


Fig. 4 Cumulative emissions of CO₂ (Pg CO₂), CH₄ (Tg CH₄) and N₂O (Tg N₂O) and their sum (3-GHG; Pg CO₂-e₁₀₀) during 1851–2021, shown globally and for selected regions. Emissions of CH₄ and N₂O are converted to units Pg CO₂-e₁₀₀ using the GWP* approach³⁸ and summed to a 3-GHG total (bottom row). The ISO3 codes of countries in the legend are: CHN, China; RUS, Russia; BRA, Brazil; IND, India; IDN, Indonesia. All data shown are provided in the **Data Records**⁴⁴.

Here, we use Eq. (2) to estimate the cumulative CO₂-equivalent ($E_{CO_2-e_{100}}$) emissions of N₂O (an LULUCF) for years 1851–2021, adopting a GWP₁₀₀ value for N₂O of 273 as reported in IPCC AR6^{2,39}. We substituted the result ($E_{CO_2-e_{100}}$) into Eq. (1) in place of E_{CO_2} to estimate the GMST response to emissions of N₂O during 1851–2021 (since the base year 1850). We adopted the TCRE value of 0.45 °C per 10³ Pg CO₂ in Eq. (1), as discussed above (see section ‘**Global Mean Surface Temperature Response to Cumulative CO₂ Emissions**’).

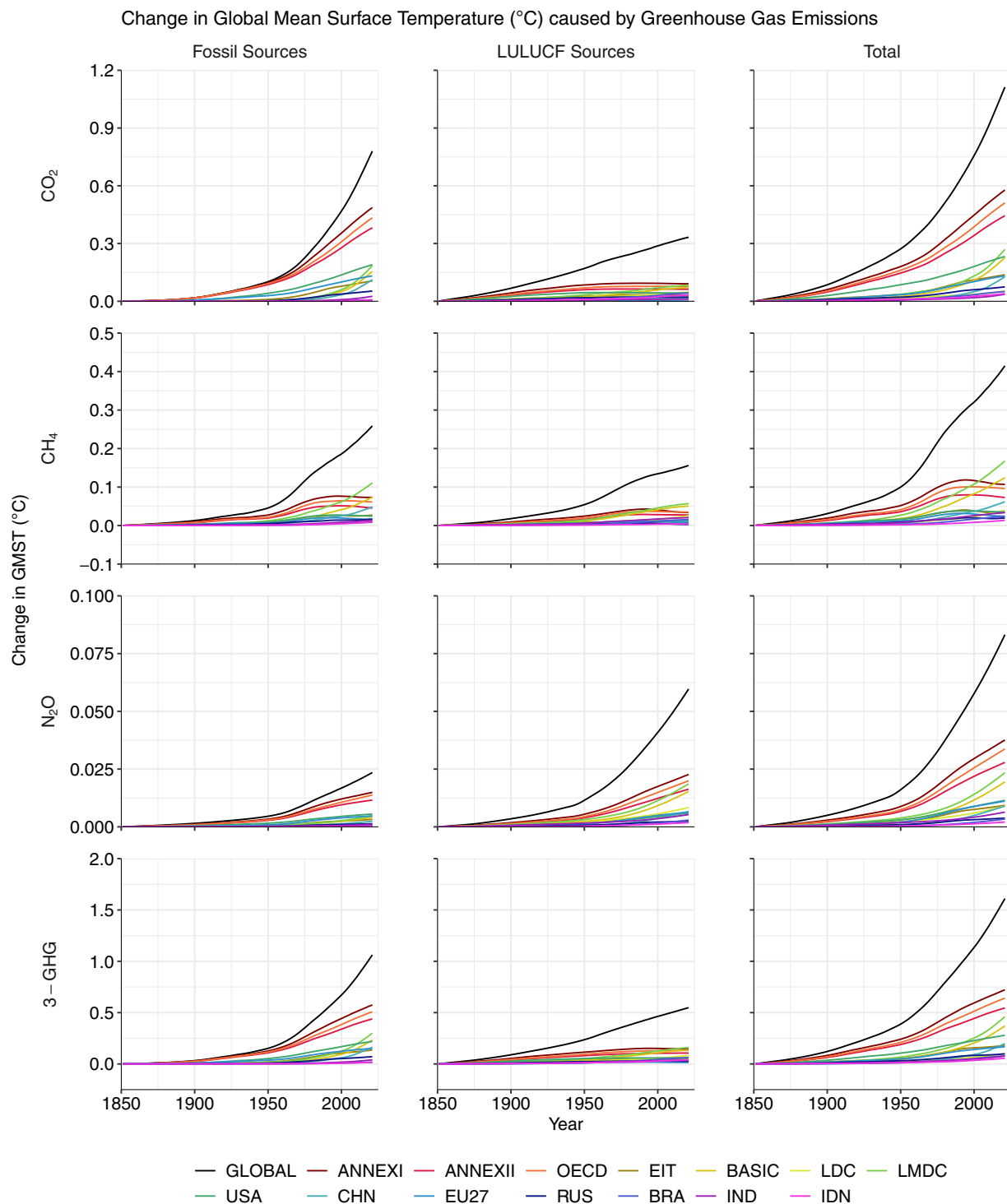


Fig. 5 Response of global mean surface temperature (GMST, °C) to emissions of CO₂, CH₄ and N₂O during 1851–2021, shown globally and for selected regions. The GMST response to historical CH₄ emissions began to fall in some developed regions towards the end of the 20th Century due to a decline in CH₄ emissions versus prior decades (see Fig. 3). The temperature response is calculated by multiplying the transient response to cumulative emissions of CO₂ (TCRE) by the total emissions of CO₂, CH₄ and N₂O expressed in Pg CO₂-e₁₀₀. Emissions of CH₄ and N₂O are converted to units Pg CO₂-e₁₀₀ using the GWP* approach³⁸. The ISO3 codes of countries in the legend are: CHN, China; RUS, Russia; BRA, Brazil; IND, India; IDN, Indonesia. All data shown are provided in the **Data Records**⁴⁴.

Non-CO₂ short-lived climate forcers (CH₄). Recent studies have highlighted that the nature of the GMST response to emissions of short-lived climate forcers (SLCFs), including CH₄, differs considerably from the GMST response to LLCFs^{39,72,73}. Due to the short atmospheric lifetime of SLCFs (e.g. ~9 years for CH₄),

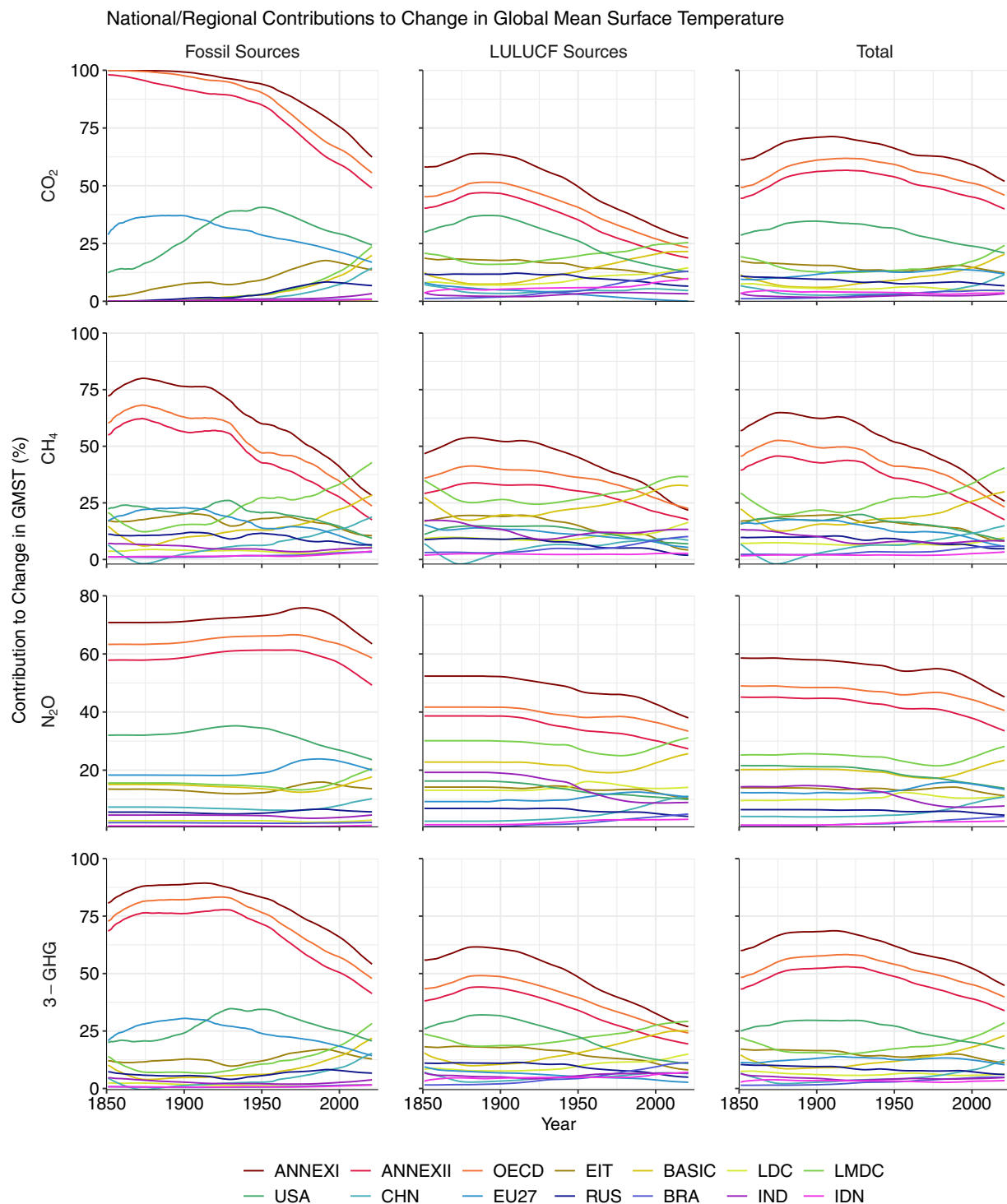


Fig. 6 National or regional contributions (%) to change in global mean surface temperature (GMST) during 1851–2021, shown for selected regions. The ISO3 codes of countries in the legend are: CHN, China; RUS, Russia; BRA, Brazil; IND, India; IDN, Indonesia. All data shown are provided in the **Data Records**⁴⁴.

the atmospheric concentration of a SLCF re-equilibrates within a short period following an increase or decrease in annual emissions. Consequently, the radiative forcing also stabilises well within a 100-year time horizon and the GMST response to SLCF emissions is not simply proportional to cumulative emissions as in the case of LLCFs. Rather, the GMST response to historical CH_4 depends foremost on recent changes in the rate of annual emissions and to a lesser extent on cumulative longer-term emissions^{36,38,39}.

To account for the differing dynamics of SLCFs in the atmosphere over long time horizons, recent work has focussed on providing an adaptation to Eq. (2) that includes the response of GMST to both cumulative SLCF

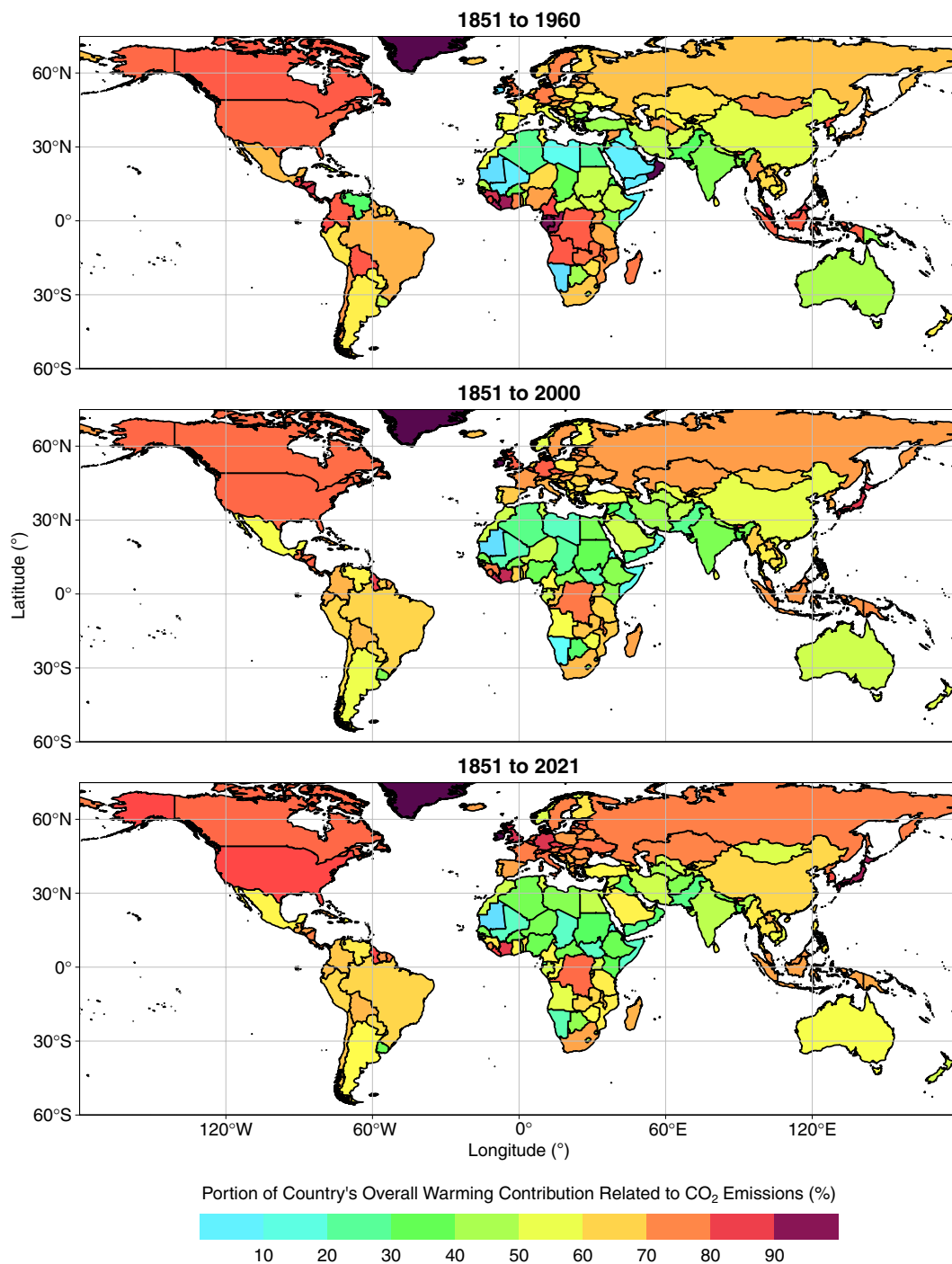


Fig. 7 Percentage of each country's total contribution to change in global mean surface temperature (GMST) related to emissions of CO₂ (as opposed to CH₄ or N₂O), over three time periods. All data shown are provided in the **Data Records**⁴⁴.

emissions and recent changes in the rate of SLCF emissions^{35–39}. The resulting method, referred to as the GWP* approach, calculates the cumulative CO₂-equivalent emissions over a 100-year time horizon ($E_{CO_2-e_{100}}$, unit Pg CO₂-e₁₀₀) of an SLCF using Eq. (3):

$$E_{CO_2-e_{100}} = sg(E_{SL} \times GWP_{100}) + (1 - s)g \left(\frac{\Delta E_{SL(t-\Delta t)} \times GWP_{100} \times H}{\Delta t} \right) \quad (3)$$

where $\Delta E_{SL(t-\Delta t)}$ is the change in cumulative emissions (Pg SLCF) between year $t-\Delta t$ and year t , with Δt representing a recent time period (years) during which cumulative SLCF emissions have evolved (e.g. 20 years).

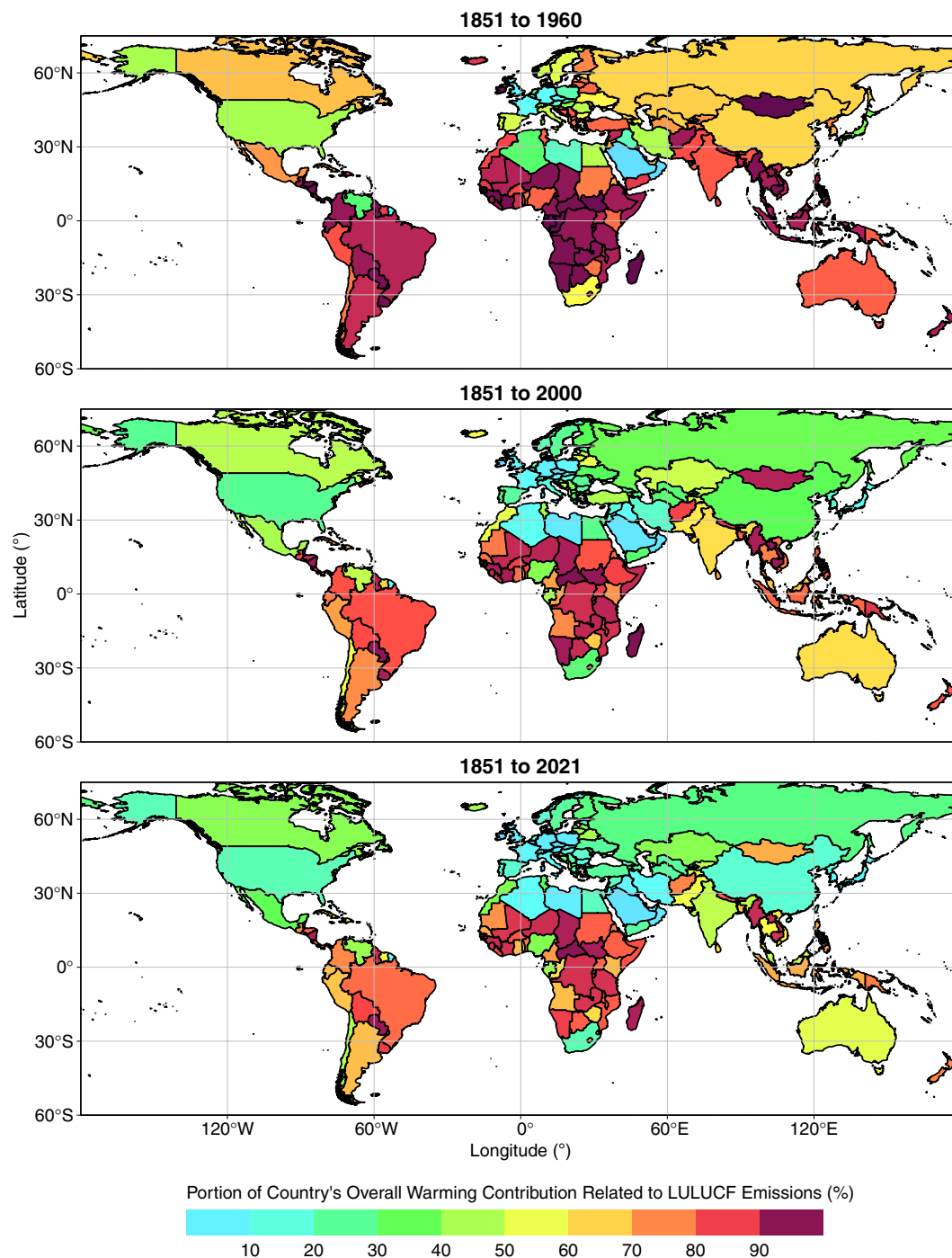


Fig. 8 Percentage of each country's total contribution to change in global mean surface temperature (GMST) related to emissions from land use, land use change and forestry (LULUCF; as opposed to fossil sources), over three time periods. These data relate to the total CO₂-equivalent emissions of CO₂, CH₄ and N₂O. All data shown are provided in the **Data Records**⁴⁴.

Note that $\frac{\Delta E_{SL}(t-\Delta t)}{\Delta t}$ is equivalent to the net change in the annual rate of SLCF emissions during the period Δt . H (years) is the time horizon of interest consistent with the time horizon of the GWP metric (e.g. 100 years is used for the GWP₁₀₀ metric). Previous studies have demonstrated the validity of Eq. (3) with $\Delta t = 20$ years and $H = 100$ (refs. ^{36,38}) and thus we also adopt these values in the current work.

The coefficient s shares a relationship with the rate of decline of radiative forcing resulting from the deep ocean thermal adjustment to recent forcing changes, ρ ($\rho = \frac{s}{H(1-s)}$) (ref. ³⁶). Based on the optimised reproduction of the GMST response given by climate models across multiple emission scenarios, $s = 0.25$ for $H = 100$ and $\rho = 0.33\% \text{ year}^{-1}$ (ref. ³⁶). The value of $\rho = 0.33\% \text{ year}^{-1}$ derives from a standard IPCC impulse-response

model^{36,74,75}. Note that the small value of s indicates that the impact of SLCF emissions on GMST is primarily determined by recent net change in the annual rate of SLCF emissions.

The coefficient g is a correction to s required to equate the radiative forcing of a SLCF directly to CO₂ forcing without reference to the temperature response³⁸. Specifically, g is a function of s ($g = \frac{(1 - e^{(-s/(1-s)))}}{s}$) and holds a value of 1.13 for $s = 0.25$ (ref. ³⁸). To calculate the response of GMST to emissions of the SLCF, $E_{CO_2-e_{100}}$ is substituted into Eq. (1) in place of E_{CO_2} :

IPCC AR6 explicitly notes that the GMST response to SLCFs such as CH₄ is more accurately reproduced when using the GWP* approach (Eq. 3) than the GWP₁₀₀ approach (Eq. 2)⁶. Treating CH₄ as a long-lived gas (i.e. calculating its CO₂ equivalent emissions using Eq. (2)) leads to overestimation of the GMST response to constant methane emissions over a multi-decadal period and a corresponding underestimation of the GMST response to any additional emissions that have been introduced over a more recent period (e.g. 20 years)^{6,37,39}.

Here, we use Eq. (3) to estimate the cumulative CO₂-equivalent emissions of CH₄ ($E_{CO_2-e_{100}}$). The GWP₁₀₀ values used in Eq. (3) were 29.8 for fossil CH₄ emissions and 27.2 for LULUCF CH₄ emissions, as reported in IPCC AR6 (difference between the two is the addition of CO₂ to the atmosphere after fossil CH₄ is oxidised)^{6,39}. In addition, we use a time horizon of $H = 100$ years and a recent time period of $\Delta t = 20$ years for assessment of the GMST response to net change in the annual rate of CH₄ emissions. For the other constants in Eq. 3, we adopt the values published by Smith *et al.* (ref. ³⁸), specifically: $s = 0.25$; $g = 1.13$. This combination of coefficient values is identical to that derived by Smith *et al.* (ref. ³⁸) and employed in IPCC AR6⁶.

For the calculation of cumulative CO₂-equivalent ($E_{CO_2-e_{100}}$) emissions of CH₄ using Eq. (3) during years 1850–1869, an emissions time series for 1830–1849 is also required (noting that $\Delta t = 20$ years). $E_{CO_2-e_{100}}$ was initially calculated for years 1850–2021 using the cumulative emissions of CH₄ (E_{SL}) since 1850 and net changes in the annual rate of SLCF emissions ($\frac{\Delta E_{SL(t-\Delta t)}}{\Delta t}$) prior to 1850 are also included where necessary. Thereafter, the estimates of $E_{CO_2-e_{100}}$ for years 1851–2021 were re-based to the year 1850 by subtracting the value of $E_{CO_2-e_{100}}$ in year 1850 from its value in each year 1851–2021.

We substituted the resulting value of $E_{CO_2-e_{100}}$ into Eq. (1) in place of E_{CO_2} to estimate the GMST response to emissions of CH₄ during 1851–2021 (since the base year 1850). We adopted the TCRE value of 0.45 °C per 10³ Pg CO₂ in Eq. (1) as discussed above (see section ‘Global Mean Surface Temperature Response to Cumulative CO₂ Emissions’).

National and international contributions to warming. The express Eqs. 1–3 were applied as described above to the global emissions records *and* to the records of emissions for individual countries and country groupings, *and* with subdivisions for fossil and LULUCF emissions. Thereafter, the fractional contributions of each country to change in GMST was estimated by dividing the GMST response to national emissions by the GMST response to global emissions. Note that the contributions of each sector (fossil and LULUCF) to change in GMST sum linearly to the national contributions to change in GMST, while the national contributions to change in GMST sum linearly to the global change in GMST. Hence the estimates of GMST response from our express approach are explicitly additive, allowing decomposition of warming contributions to countries and sectors.

The contributions of various country groupings to emissions and the GMST response to emissions were calculated by summing the contributions of constituent nations. The country groupings considered in this study were as follows. Several groupings derive from UNFCCC definitions (<https://unfccc.int/parties-observers>), including the 42 Annex I parties, the 23 Annex II parties (the most economically developed members of the Annex I), the 15 Economies in Transition (EIT; the lesser-developed members of Annex I), the 47 Least Developed Countries (LDCs), and the group of 24 Like-Minded Developing Countries (LMDC). In addition, we consider the contributions of the group of 38 countries of the Organisation for Economic Co-operation and Development (OECD; <https://www.oecd.org/about/document/ratification-oecd-convention.htm>), the group of 27 members of the European Union (EU27; https://european-union.europa.eu/principles-countries-history/country-profiles_en), and the Brazil, South Africa, India and China (BASIC) group. Lists of the countries included in each country grouping are provided in the **Data Records**⁴⁴.

Uncertainty assessment. Here, we characterise the uncertainties for all terms in Eqs. (1–3) at a consistent 1-sigma level (68% confidence interval), thus enabling the propagation of errors in related work if desired. We do not provide explicit estimates of uncertainty within our data records⁴⁴.

Uncertainty in emissions estimates. Our estimates of CO₂ emissions derive from the global carbon budget of the GCP^{1,32,33}. The GCP provides an expert judgement of the uncertainty in its CO₂ emissions estimates. At the global scale and in Annex I countries reporting to the UNFCCC, the GCP estimates that 1-sigma uncertainties in annual fossil CO₂ emissions are 5%^{3,40,64}. For non-Annex countries the GCP estimates a 10% uncertainty in annual fossil CO₂ emissions due to less stringent reporting and verification. Meanwhile, the GCP estimates the 1-sigma uncertainty in global LULUCF CO₂ emissions to be 50%³. The uncertainty on national scales is poorly constrained but likely higher than 50%³. We note that data relating to LULUCF emissions in China are subject to considerable uncertainty because the LUH2 and FRA datasets show opposing signs⁵¹. FRA exhibits large-scale forest plantation in China since the mid 20th Century, leading to an LULUCF sink, whereas LUH2 indicates widespread forest loss in China over the same period⁵¹.

As discussed above, Minx *et al.* (ref. ⁶⁴) compared available estimates of CH₄ and N₂O emissions and found that PRIMAP-hist (TP scenario) lies centrally amongst those estimates. The spread of the estimates may be partly indicative of uncertainty in CH₄ and N₂O, although commonalities in parameter choice and requisite

data sources amongst the emissions datasets means that uncertainties are likely to be larger than inferred by the spread of estimates alone⁴⁴. Minx *et al.* (refs. 58,64) provide a best judgement of 1-sigma uncertainty in total CH₄ emissions during 1970–2018 of $\pm 30\%$ globally, with higher uncertainties nationally and for earlier decades⁴⁴. Their current best judgement of 1-sigma uncertainty in total N₂O emissions during 1970–2018 is $\pm 60\%$ globally, and higher nationally and for earlier decades⁶⁴. Hence, we note that the CH₄ and N₂O emissions estimates used in the current study lie centrally within a large uncertainty range globally and within a poorly constrained uncertainty range on national scales.

Uncertainty in the transient climate response to cumulative CO₂ emissions. The IPCC AR6^{2,6} TCRE (k in Eq. (1)) considered 27 assessments of the TCRE value published between 2009 and 2021 and revised the earlier estimate provided in AR5 based on 17 subsequent estimates. Each estimate involved a different model or model ensemble exhibiting a particular climate sensitivity, with some studies constraining their estimates to observed changes in surface temperature. The revised best estimate for TCRE was 1.65 °C per 10³ Pg C emitted (0.45 °C per 10³ Pg CO₂ emitted)^{4,6} with a 90% confidence interval of 1.0–2.3 °C per 10³ Pg C emitted, corresponding to 1-sigma uncertainty of 0.4 °C per 10³ Pg C emitted (0.18 °C per 10³ Pg CO₂ emitted).

Uncertainty in the global warming potential of CH₄ and N₂O. The IPCC AR6² reviewed the current understanding of uncertainty in the global warming potential of CH₄ and N₂O and many other anthropogenic gases and aerosols. Based on the latest evidence, the report arrives at a current best estimate of 29.8 for the GWP₁₀₀ of fossil CH₄, with an uncertainty of ± 11.9 at the 90% confidence interval, corresponding to a 1-sigma uncertainty of ± 7.1 . The GWP₁₀₀ of biogenic CH₄ is reported to be 27.2 with an uncertainty of ± 10.9 at the 90% confidence interval, corresponding to 1-sigma uncertainty of ± 6.5 . The major sources of uncertainty that contribute to the overall uncertainty in the GWP₁₀₀ of CH₄ are (in order of magnitude): (i) the absolute global warming potential of CO₂ (the time-integrated radiative forcing caused by a CO₂ emissions pulse over the 100-year horizon); (ii) the measurement of radiative efficiency of CH₄ (absorption of energy across a spectrum of wavelengths); (iii) the atmospheric lifetime of the gas, and; (iv) chemistry feedbacks (interactions with other atmospheric gases). The GWP₁₀₀ of N₂O is reported to be 273 with an uncertainty of ± 128.3 at the 90% confidence interval, corresponding to 1-sigma uncertainty of ± 76.4 . The major sources of uncertainty that contribute to the overall uncertainty in the GWP₁₀₀ of N₂O differ from those of CH₄ and are (in order of magnitude): (i) chemistry feedbacks; (ii) the absolute global warming potential of CO₂, and; (iii) the measurement of radiative efficiency of N₂O.

Other uncertainties. The coefficient s relates to the rate of decline of radiative forcing resulting from the deep ocean thermal adjustment to recent forcing changes, ρ (see **Methods**). The value of ρ used here (0.33% year⁻¹) reflects a multi-model average deep ocean adjustment period of around 300 years ($1/\rho$). The deep ocean adjustment period depends on the climate model used to derive it, with a 1-sigma range of around ± 110 years^{36,76}. The uncertainty range of the deep ocean adjustment period infers an uncertainty in ρ of 0.24–0.51% year⁻¹ at the 1-sigma interval. Finally, g in Eq. (3) is a function of s (see **Methods**) such that its uncertainty corresponds to that of s .

Data Records

All records are available via a Zenodo data repository (ref. 44). The data records include three comma separated values (.csv) files as described below. All files are in 'long' format with one value provided in the *Data* column for each combination of the categorical variables *Year*, *Country Name*, *Country ISO3 code*, *Gas*, and *Component* columns. The *Component* field specifies fossil emissions, LULUCF emissions or total emissions of the gas. *Gas* specifies CO₂, CH₄, N₂O or the three-gas total (labelled 3-GHG). ISO3 codes are specifically the unique ISO 3166-1 alpha-3 codes of each country (<https://www.iso.org/iso-3166-country-codes.html>).

EMISSIONS_ANNUAL_1830–2021.csv (26.3 MB). *Data* includes annual emissions of CO₂ (Pg CO₂ year⁻¹), CH₄ (Tg CH₄ year⁻¹) and N₂O (Tg N₂O year⁻¹) during 1830–2021. The *Data* column provides values for every combination of the categorical variables. There are 369,048 data rows in the current version. Note that data for the years 1830–1849 are provided as these data are needed for the calculation of CO₂-equivalent emissions of CH₄ using the GWP* approach during years 1850–1869 (see Eq. 3).

EMISSIONS_CUMULATIVE_CO2e100_1851–2021.csv (33.8 MB). *Data* includes the cumulative CO₂ equivalent emissions in units Pg CO₂-e₁₀₀ during 1851–2021. The *Data* column provides values for every combination of the categorical variables. There are 450,585 data rows in the current version.

GMST_response_1851–2021.csv (28.9 MB). *Data* includes the change in global mean surface temperature (GMST) due to emissions of the three gases during 1851–2021 in units °C. The *Data* column provides values for every combination of the categorical variables. There are 450,585 data rows in the current version.

In addition to the data records above, we provide lists of the countries included in each of the following country groupings in a Microsoft Excel workbook named **COUNTY_GROUPINGS.xlsx (~20 KB)**: the 42 Annex I parties; the 23 Annex II parties; the 15 Economies in Transition (EIT); the 47 Least Developed Countries (LDCs); the 24 Like-Minded Developing Countries (LMDC); the 38 countries of the Organisation for Economic Co-operation and Development (OECD); the 27 members of the European Union (EU27), and the four countries of the BASIC group. Each group occupies one worksheet of the Excel workbook and consists of one column of listed countries.

Technical Validation

We provide Figs. 1–8, Tables 1–4, and the text below to assist with the technical validation of our dataset including its cross-comparison with other studies.

Cumulative emissions. According to our estimates, 2,471 Pg CO₂ were emitted globally during 1851–2021 (Figs. 3, 4) leading to a CO₂-induced increase in GMST of 1.11 °C (Fig. 5, Table 1). We estimate that cumulative fossil CO₂ emissions during 1851–2021 amounted to 1,732 Pg CO₂ (Figs. 3, 4) and caused warming of 0.78 °C (Fig. 5, Table 1). The dataset of fossil CO₂ emissions is taken directly from the GCP⁴⁰ such that the cumulative emissions for 1851–2021 match that reported in the GCB (ref. ³). Our estimate of cumulative LULUCF CO₂ emissions during 1851–2021 is 739 Pg CO₂ (Figs. 3, 4) leading to warming of 0.33 °C (Fig. 5, Table 1). The estimate of global cumulative LULUCF CO₂ emissions also matches the value reported in the GCB³ because the GCB also uses the average estimate from the three products employed here^{49–51}. Thus, our estimates of total annual and cumulative CO₂ emissions match the estimates made by the GCP on the global scale³.

According to the PRIMAP-hist HISTTP dataset, 921 Pg CO₂-e₁₀₀ of CH₄ (25,669 Tg CH₄) were emitted globally during 1851–2021 (Figs. 3, 4) leading to a 0.41 °C increase in GMST (Fig. 5, Table 2). We estimate that cumulative fossil CH₄ emissions during 1851–2021 amounted to 575 Pg CO₂-e₁₀₀ (12,803 Tg CH₄, Figs. 3, 4) and caused warming of 0.26 °C (Fig. 5, Table 2), while cumulative LULUCF CH₄ emissions amounted to 347 Pg CO₂-e₁₀₀ (12,866 Tg CH₄, Figs. 3, 4) and caused warming of 0.16 °C (Fig. 5, Table 2). Also relevant to the calculation of CO₂-equivalent emission of CH₄ are the net changes in the rate of annual emissions over the past 20 years ($\frac{\Delta E_{SL}(t-\Delta t)}{\Delta t}$ in Eq. (3)), which were 334 Tg CH₄ year⁻¹ during 2002–2021 including 198 Tg CH₄ year⁻¹ from fossil sources and 135 Tg CH₄ year⁻¹ from LULUCF sources (Fig. 3). Our CH₄ emissions estimates are taken directly from the PRIMAP-hist dataset such that the cumulative emissions for 1851–2021 match the PRIMAP-hist record^{41,42}.

According to the PRIMAP-hist HISTTP dataset, 185 Pg CO₂-e₁₀₀ of N₂O (677 Tg N₂O, Figs. 3, 4) were emitted globally during 1851–2021 leading to a 0.08 °C increase in GMST (Fig. 5, Table 3). We estimate that cumulative fossil N₂O emissions during 1851–2021 amounted to 52 Pg CO₂-e₁₀₀ (191 Tg N₂O) and caused warming of 0.02 °C (Fig. 5, Table 3), while cumulative LULUCF N₂O emissions amounted to 133 Pg CO₂-e₁₀₀ (485 Tg N₂O, Figs. 3, 4) and caused warming of 0.06 °C (Fig. 5, Table 3). Our N₂O emissions estimates are taken directly from the PRIMAP-hist dataset such that the cumulative emissions for 1851–2021 match the PRIMAP-hist record^{41,42}.

Minx *et al.* (ref. ⁶⁴) recently compared seven estimates of total CH₄ and N₂O emissions for the years 1970–2018 and found that the PRIMAP-hist dataset (with the HISTTP scenario used here) yields estimates that lie centrally within the range of available estimates. For example, the PRIMAP-hist estimate for total fossil CH₄ emissions in 2010 is around 340 Tg CH₄ year⁻¹, and there are three higher estimates (up to 360 Tg CH₄ year⁻¹) and three lower estimates (lowest estimate 300 Tg CH₄ year⁻¹). Also for the year 2010, the PRIMAP-hist estimate for total fossil N₂O emissions in 2010 is around 9.5 Tg N₂O year⁻¹, and there are two higher estimates (up to 12 Tg N₂O year⁻¹) and four lower estimates (lowest estimate 8.5 Tg N₂O year⁻¹). PRIMAP-hist similarly lies centrally amongst estimates of CH₄ and N₂O emissions in the years 1970, 1980, 1990 and 2018 (see Fig. 1c,d of ref. ⁶⁴).

Change in global mean surface temperature. We compared our estimates of change in GMST (Fig. 5) caused by historical emissions of CO₂, CH₄, and N₂O with the values reported recently by the IPCC in AR6^{1,8}. The IPCC AR6 estimated the change in GMST caused by CO₂, CH₄ and N₂O emissions to be 1.4 °C between 1850–1900 and 2010–2019^{1,8}. For a close comparison, we calculated the change in GMST between the periods 1851–1900 and 2010–2019 to be 1.39 °C in our dataset.

The IPCC AR6 estimated the change in GMST caused by cumulative CO₂ emissions from all sources to be 0.8 °C between 1850–1900 and 2010–2019, with an uncertainty range of 0.5–1.2 °C (90% confidence interval)^{1,8}. For a close comparison, we calculated the change in GMST between the periods 1851–1900 and 2010–2019 from our dataset to be 0.96 °C. Hence, our estimate of warming caused by global historical CO₂ emissions lies within the very likely range identified by the IPCC, and 0.16 °C above the IPCC's central estimate.

The IPCC AR6 estimated the change in GMST due to historical CH₄ emissions from all sources to be 0.5 °C between 1850–1900 and 2010–2019, with an uncertainty range of 0.3–0.8 °C (90% confidence interval)^{1,8}. For a close comparison, we calculated the change in GMST between 1851–1900 and 2010–2019 to be 0.37 °C in our dataset. Hence, our estimate of warming caused by global historical CH₄ emissions lies within the very likely range identified by the IPCC, and 0.13 °C below the central estimate.

The IPCC AR6 estimated the change in GMST due to historical N₂O emissions from all sources to be 0.09 °C between 1850–1900 and 2010–2019, with an uncertainty range of 0.05–0.16 °C (90% confidence interval)^{1,8}. For a close comparison, we calculated the change in GMST between 1851–1900 and 2010–2019 to be 0.07 °C in our dataset. Hence, our estimate of warming caused by global historical N₂O emissions lies within the very likely range identified by the IPCC, and 0.02 °C below the central estimate.

National contributions to warming. *Contributions through CO₂ emissions.* The ranking of countries by national contributions to warming through CO₂ emissions presented here compares well with those published previously. For example, the list of top-10 contributors to the warming caused by total CO₂ emissions during 1851–2021 (USA, China, Russia, Brazil, Germany, Indonesia, India, UK, Japan and Canada; Table 1, Figs. 5, 6) shares many similarities with the list presented by Matthews *et al.* (ref. ¹² USA, China, Russia, Brazil, India, Germany, UK, France, Indonesia and Canada) for the years 1800–2005. Skeie *et al.*¹² also arrived at a similar list (USA, EU28, China, Russia, Indonesia, Brazil, Japan, India, Canada) for the years 1850–2012 (note that the EU27

post-Brexit also ranks between the USA and China in our dataset; Table 1, Figs. 5, 6). Some differences in rank between studies may be explained by differences in the period of emissions considered. For example, our estimates indicate that China moved into second position ahead of Russia since 2005 based on its evolving emissions since 1850, while India rose from tenth position to fifth position ahead of Indonesia, Germany, the UK, Japan and Canada. Other differences versus previous work may result from revisions to emissions estimates, and we particularly note that the order of national contributions to warming caused by LULUCF CO₂ emissions shared fewer similarities with that of previous work than in the case of warming related to fossil CO₂ emissions. This broadly aligns with expectations as LULUCF CO₂ emissions estimates carry greater uncertainty and are subject to more frequent and substantial revisions than fossil CO₂ emissions estimates^{3,77}.

Overall contributions to warming through total CO₂ emissions mask large differences in the relative contributions of fossil and LULUCF emissions at national levels (Table 1, Fig. 8). For example, Brazil's contribution to warming from LULUCF CO₂ emissions (0.04 °C, Table 1, Figs. 5, 6) is greater than any other nation's contribution and accounts for most of its overall contribution to warming through total CO₂ emissions (0.05 °C, Table 1, Figs. 5, 6). Indonesia's contributions to warming through historical CO₂ emissions are also dominated by LULUCF emissions. These observations are in line with prior reports of a strong contribution of LULUCF emissions to total CO₂ emissions in Brazil and Indonesia⁷⁸. These examples highlight how excluding LULUCF emissions can lead to a substantial underestimation of the contribution of some nations to warming^{12,13,16}. In stark contrast, LULUCF emissions during 1851–2021 were negative in several European countries (including Germany and France) and thus the LULUCF sector in these countries contributed to a cooling of GMST, slightly offsetting the warming associated with their fossil CO₂ emissions. The cooling effect results from negative cumulative emissions from the LULUCF sector of these countries since 1850 as observed by Friedlingstein *et al.* (ref. 3). This pattern may in part reflect unaccounted LULUCF emissions of CO₂ that occurred as a result of land use changes in Europe prior to 1850, although previous work has established that accounting for preindustrial LULUCF prior to 1850 impacts the share a region takes of global warming only by a few percent and is thus of similar magnitude as uncertainties related to other methodological choices (for the region itself, however, its contribution may be altered substantially in relative terms)^{16,79}.

Globally, 30% of the warming caused by total CO₂ emissions during 1851–2021 is associated with LULUCF, but with largely varying shares in different countries (Fig. 8). The contribution of LULUCF CO₂-related warming to total CO₂-related warming lies around the global value in Russia and India, whereas the LULUCF share of CO₂-related warming is lower in the US (18%) and China (12%) and higher in Brazil (85%) and Indonesia (83%). The cooling related to LULUCF CO₂ emissions in Germany and France offsets 3–6% of the warming related to their fossil CO₂ emissions. While the developed members of the OECD and Annex II contribute more towards fossil CO₂-related warming than the like-minded developing countries (LMDC) and non-Annex groups, the opposite is true with respect to LULUCF CO₂-related warming.

Our focus here has been on contributions to warming during the period 1851–2021, however we note that national contributions have evolved through time since 1850 (Figs. 5, 6). For example, the contributions of the LMDC group to LULUCF CO₂-related only overtook the contributions of the OECD around the year 2011.

Effect of including methane and nitrous oxide. Major shares of global CH₄ and N₂O emissions are associated with land use (Tables 2, 3; Figs. 3, 4)^{4,5}. The GCP has estimated that livestock (enteric fermentation and manure management) and rice cultivation contributed almost 40% of global total CH₄ emission during 2009–2018⁴, and that nitrogen fertilizer use, livestock and manure management contribute 50% of global total N₂O emissions during 2007–2016⁵. Consequently, previous studies have found that considering emissions of CH₄ and N₂O tends to increase the contributions to warming of countries with high agricultural intensity and area under management^{12,13}. We observe similar patterns in our dataset (Tables 2, 3; Figs. 5–8). Globally, the LULUCF sector accounted for 38% of the total warming from CH₄ emissions and 72% of the warming from N₂O emissions during 1851–2021 (Tables 2, 3; Figs. 5, 6). Notably, a considerably greater fraction of the warming caused by CH₄ emissions was associated with LULUCF in some countries (e.g. over 60% in India, Brazil, Australia and Pakistan; Table 2; Figs. 5, 6). When considering CH₄ and N₂O emissions, the contribution to warming of India, China, and Brazil rose by 110%, 56% and 55%, respectively, relative to the CO₂-related warming alone (Tables 1, 4; Figs. 5–7). For comparison, the additional contribution to warming of most other large emitters (USA, EU27, Russia, Canada) was 30% or lower, and values were below 15% in Germany, the UK and Japan.

The differential effects of including CH₄ and N₂O in assessments of national contributions to warming leads to a re-ordering of the top contributors (Table 4) as compared to the scenario in which only CO₂-related warming is considered (Table 1; Figs. 5, 6). For example, India moves from 7th to 5th position above Indonesia and Germany, and China's contribution also moves beyond that of the EU27. Hence we highlight the critical importance of study design when ranking the contributions of individual nations to warming in line with previous studies^{12,13}. Moreover, we note that emissions of CH₄ and N₂O are more uncertain than emissions of fossil CO₂ and so any changes in the ranking of contributors related to the additional consideration of CH₄ and N₂O should be treated with particular scrutiny.

Due to re-equilibration of atmospheric CH₄ concentrations within a short period after a change in annual emissions, it is possible for a reduction in the rate of annual CH₄ emissions to bring about a cooling effect on GMST even if the annual emissions remain positive, unlike in the case of CO₂ or N₂O emissions. For an example of this effect, see Figs. 5, 6 where China's contribution to change in GMST is negative (signifying a cooling effect) during years 1866–1885. China's mean annual CH₄ emissions fell by around 8% between the 1850s and 1870s according to the PRIMAP-hist dataset (Fig. 3), resulting in negative cumulative CO₂-equivalent emissions in line with Eq. (3) and a small cooling effect on GMST until the year 1885 (Figs. 5, 6). Similarly, emissions of CH₄ began to decline in many developed nations in the late decades of the 20th Century, resulting in a reduced

Gas	Source	Country or Grouping	°C	%
CO ₂	Total	GLOBAL	1.11	
		ANNEXI	0.58	51.9
		OECD	0.51	45.9
		NONANNEX	0.51	45.8
		ANNEXII	0.44	40.0
		LMDC	0.27	24.2
		USA	0.23	20.9
		BASIC	0.23	20.4
		EIT	0.14	12.4
		EU27	0.13	11.8
		China	0.13	11.5
		Russia	0.07	6.7
		LDC	0.05	4.6
		Brazil	0.05	4.6
		Germany	0.04	3.7
		Indonesia	0.04	3.6
		India	0.04	3.3
		United Kingdom	0.03	3.1
		Fossil	GLOBAL	0.78
	ANNEXI		0.49	62.4
	OECD		0.43	55.6
	ANNEXII		0.38	49.0
	NONANNEX		0.27	34.3
	USA		0.19	24.4
	LMDC		0.18	23.7
	BASIC		0.16	19.9
	EU27		0.13	16.9
	China		0.11	14.4
	EIT		0.11	13.6
	Russia		0.05	6.8
	Germany		0.04	5.4
	United Kingdom		0.03	4.3
	Japan		0.03	3.9
	India		0.03	3.3
	LULUC	GLOBAL	0.33	
		NONANNEX	0.24	72.6
		ANNEXI	0.09	27.3
		LMDC	0.08	25.4
		OECD	0.08	23.3
		BASIC	0.07	21.5
		ANNEXII	0.06	18.8
		LDC	0.05	14.5
Brazil		0.04	13.0	
USA		0.04	12.8	
Indonesia		0.03	9.9	
EIT		0.03	9.6	
Russia		0.02	6.5	
China		0.02	4.6	
Canada		0.01	3.4	
India		0.01	3.2	

Table 1. Change in global mean surface temperature (GMST) due to emissions of CO₂ during 1851–2021, and the contributions of countries or country groupings which contributed at least 3% of the change in GMST. Bold text is used to highlight individual countries as opposed to country groupings, while bold italicised text is used to highlight the European Union as its 27 constituent countries share common environmental, agricultural and climate policies. The **Data Records**⁴⁴ include all data shown, as well as the values for all countries and country groupings.

Gas	Source	Country or Grouping	°C	%
CH ₄	Total	GLOBAL	0.41	
		NONANNEX	0.31	74.0
		LMDC	0.17	40.5
		BASIC	0.12	29.9
		ANNEXI	0.11	25.8
		OECD	0.10	23.2
		ANNEXII	0.07	17.5
		China	0.06	14.9
		LDC	0.04	9.6
		USA	0.04	8.5
		India	0.03	8.2
		EIT	0.03	8.1
		Brazil	0.02	5.9
		EU27	0.02	5.8
		Russia	0.02	4.8
	Indonesia	0.01	3.3	
	Fossil	GLOBAL	0.26	
		NONANNEX	0.18	71.5
		LMDC	0.11	42.8
		BASIC	0.07	28.4
		ANNEXI	0.07	28.2
		OECD	0.06	23.6
		China	0.05	18.8
		ANNEXII	0.05	17.5
		EIT	0.03	10.5
		USA	0.02	9.4
		Russia	0.02	6.5
		EU27	0.02	6.0
		LDC	0.01	5.5
		India	0.01	5.1
		Indonesia	0.01	3.7
	Nigeria	0.01	3.7	
	Brazil	0.01	3.3	
	LULUC	GLOBAL	0.16	
		NONANNEX	0.12	78.2
		LMDC	0.06	36.6
		BASIC	0.05	32.5
		OECD	0.04	22.5
		ANNEXI	0.03	21.7
		ANNEXII	0.03	17.6
		LDC	0.03	16.4
		India	0.02	13.3
Brazil		0.02	10.1	
China		0.01	8.5	
USA		0.01	6.9	
<i>EU27</i>	<i>0.01</i>	5.3		
EIT	0.01	4.2		
Pakistan	0.00	3.2		
Australia	0.00	3.1		

Table 2. Change in global mean surface temperature (GMST) due to emissions of CH₄ during 1851–2021, and the contributions of countries or country groupings which contributed at least 3% of the change in GMST. Bold text is used to highlight individual countries as opposed to country groupings, while bold italicised text is used to highlight the European Union as its 27 constituent countries share common environmental, agricultural and climate policies. The **Data Records**⁴⁴ include all data shown, as well as the values for all countries and country groupings.

Gas	Source	Country or Grouping	°C	%	
N ₂ O	Total	GLOBAL	0.083		
		NONANNEX	0.045	54.7	
		ANNEXI	0.038	45.2	
		OECD	0.034	40.5	
		ANNEXII	0.028	33.5	
		LMDC	0.023	28.2	
		BASIC	0.019	23.4	
		USA	0.011	13.8	
		<i>EU27</i>	<i>0.011</i>	<i>13.4</i>	
		EIT	0.009	11.2	
		LDC	0.009	10.9	
		China	0.009	10.8	
		India	0.006	7.7	
		Russia	0.004	4.5	
		Brazil	0.003	4.1	
		Australia	0.003	3.8	
		Fossil	GLOBAL	0.024	
			ANNEXI	0.015	63.5
	OECD		0.014	58.6	
	ANNEXII		0.012	49.2	
	NONANNEX		0.009	36.2	
	USA		0.006	23.6	
	LMDC		0.005	20.6	
	<i>EU27</i>		<i>0.005</i>	<i>20.0</i>	
	BASIC		0.004	17.7	
	EIT		0.003	13.6	
	China		0.002	10.2	
	Russia		0.001	5.6	
	India		0.001	4.6	
	Germany		0.001	4.5	
	Mexico		0.001	3.6	
	France		0.001	3.6	
	United Kingdom		0.001	3.5	
	LULUC		GLOBAL	0.060	
		NONANNEX	0.037	61.9	
		ANNEXI	0.023	38.0	
		OECD	0.020	33.4	
		LMDC	0.019	31.2	
		ANNEXII	0.016	27.3	
		BASIC	0.015	25.7	
		LDC	0.008	14.1	
		China	0.007	11.0	
<i>EU27</i>		<i>0.006</i>	<i>10.8</i>		
EIT		0.006	10.2		
USA		0.006	9.9		
India		0.005	8.9		
Brazil		0.003	4.9		
Australia		0.003	4.9		
Russia		0.002	4.1		
Indonesia	0.002	3.1			

Table 3. Change in global mean surface temperature (GMST) due to emissions of N₂O during 1851–2021, and the contributions of countries or country groupings which contributed at least 3% of the change in GMST. Bold text is used to highlight individual countries as opposed to country groupings, while bold italicised text is used to highlight the European Union as its 27 constituent countries share common environmental, agricultural and climate policies. The **Data Records**⁴⁴ include all data shown, as well as the values for all countries and country groupings.

Gas	Source	Country or Grouping	°C	%	
3-GHG	Total	GLOBAL	1.61		
		NONANNEX	0.86	53.5	
		ANNEXI	0.72	44.8	
		OECD	0.64	39.8	
		ANNEXII	0.55	33.8	
		LMDC	0.46	28.6	
		BASIC	0.37	23.0	
		USA	0.28	17.3	
		China	0.20	12.3	
		EIT	0.18	11.2	
		<i>EU27</i>	<i>0.17</i>	<i>10.4</i>	
		LDC	0.10	6.2	
		Russia	0.10	6.1	
		Brazil	0.08	4.9	
		India	0.08	4.8	
		Indonesia	0.06	3.4	
		Fossil	GLOBAL	1.06	
			ANNEXI	0.57	54.1
	OECD		0.51	47.9	
	NONANNEX		0.46	43.4	
	ANNEXII		0.44	41.3	
	LMDC		0.30	28.3	
	BASIC		0.23	21.9	
	USA		0.22	20.7	
	China		0.16	15.4	
	<i>EU27</i>		<i>0.15</i>	<i>14.3</i>	
	EIT		0.14	12.8	
	Russia		0.07	6.7	
	Germany		0.05	4.3	
	India		0.04	3.8	
	United Kingdom	0.04	3.4		
	LULUC	GLOBAL	0.55		
		NONANNEX	0.40	73.0	
		LMDC	0.16	29.2	
		ANNEXI	0.15	26.9	
		BASIC	0.14	25.1	
		OECD	0.13	24.2	
		ANNEXII	0.11	19.4	
		LDC	0.08	15.0	
		Brazil	0.06	11.3	
		USA	0.06	10.8	
		EIT	0.04	8.1	
Indonesia		0.04	7.1		
India		0.04	6.7		
China		0.04	6.4		
Russia	0.03	4.9			

Table 4. Change in global mean surface temperature (GMST) due to combined emissions of CO₂, CH₄ and N₂O during 1851–2021, and the contributions of countries or country groupings which contributed at least 3% of the change in GMST. Bold text is used to highlight individual countries as opposed to country groupings, while bold italicised text is used to highlight the European Union as its 27 constituent countries share common environmental, agricultural and climate policies. The **Data Records**⁴⁴ include all data shown, as well as the values for all countries and country groupings.

contribution of these countries to warming since the mid-20th Century (Figs. 5, 6). In contrast, cooling can only result from a reduction in CO₂ and N₂O emissions if cumulative emissions become negative, owing to the linear relationship between their cumulative emissions and change in GMST (Eqs. 1, 2).

Usage Notes

In addition to the uncertainties evaluated in the **Methods**, we highlight some aspects of our study design which might affect our assessment of national contributions to warming versus alternative methodologies.

Limitations of TCRE and GWP*. While use of TCRE and GWP* to attribute historical warming has the advantages of simplicity, transparency and alignment with the latest IPCC assessment, this approach also has limitations stemming from the fact that TCRE and GWP* are first-order approximations of a complex dynamic system in which radiatively active species directly or indirectly interact. TCRE is a linear approximation of the long-term global temperature response to an emission of CO₂ that is assumed to be constant through time and independent of the emission intensity or past emissions. However, individual models typically exhibit a more complex dynamic response⁷⁴, as also illustrated by the existence of a committed warming after reaching net zero CO₂ emissions (the Zero-Emissions Commitment)⁸⁰, and by the breaking of the linear approximation in case of significant temperature overshoot⁸¹. In addition, TCRE is used indiscriminately for both fossil and LULUCF sources, whereas LULUCF CO₂ emissions are known not to be precisely equivalent to fossil ones⁸² because they are mostly caused by land cover change that simultaneously reduces the land carbon sink⁸³, thereby changing the TCRE itself. In the GCB³, this is termed the loss of additional sink capacity^{3,51,84}, and accounting for it would slightly increase the relative contribution of countries having emitted significant amounts of LULUCF CO₂. Also, LULUCF exerts an associated biophysical effect on climate via changes in planetary surface albedo that is not paralleled in activities leading to fossil CO₂ emission. The biophysical effect causes differences in the climatic response to LULUCF versus fossil CO₂ emission, however this distinction is not accounted for in the TCRE value⁸⁵.

GWP* is a variation of the classic GWP that enables definition of CO₂-equivalence within the TCRE framework. It remedies what is perhaps the most critical shortcoming of GWP when applied to short-lived species: a lack of explicit time dynamic. The GWP* dynamic remains simplistic, however, compared to what would be obtained with non-linear models with a detailed evolution of a species' atmospheric lifetimes⁸⁶. Furthermore, GWP* uses GWP in its formulation, and therefore includes only the atmo- and biogeo-chemical feedbacks accounted for within the chosen GWP value. For instance, our chosen GWP for CH₄ includes effects on tropospheric ozone and the carbon cycle through the climate feedback^{2,87}, and it has implicit backgrounds of atmospheric CH₄ and emission of ozone precursors². In GWP and GWP*, these factors and feedbacks are linearised, assumed constant in time, and attributed to the main species of interest, which essentially ignores the complex real-life cross-species dynamics^{2,88}. Accounting for these, however, requires more advanced models^{89–91} that come at the cost of the simplicity and transparency we were aiming at, for changes in national contributions that would likely be of second-order.

Other considerations related to study design. Many foregoing studies have highlighted the importance of perspective when assigning contributions to climate change, specifically referring to the role of study design in determining contributions^{13,16,18,21,22,28,92–97}. Our results similarly point to various structural elements of study design that influence the assessed contributions to warming, which we summarise here.

First, contributions to warming depend on the gases and aerosols considered in the analysis. Different anthropogenic activities emit various gases and aerosols at ranging intensities (e.g. industrial versus agricultural). Each country has a unique environmental and socioeconomic situation causing differences in the prevalence of source activities and influencing emission rates of associated gases and aerosols. Consequently, a country's contribution to warming increases if a gas or aerosol associated with one of its prevalent activities is considered in the assessment. For example, the inclusion of CH₄ and N₂O enhances the contribution to warming of countries with intensive or extensive agriculture^{12,13}. Here, we consider only CO₂, CH₄ and N₂O emissions in our assessment of national contributions to warming, thus excluding national contributions to warming through emissions of other radiatively active species. The IPCC AR6 finds that anthropogenic emissions of black carbon aerosols, halogenated gases (CFC + HCFC + HFC) and volatile organic compounds and carbon monoxide (NMVOC + CO) cause a warming at the global scale comparable to that of N₂O^{1,8}. Inclusion of these species thus has potential to influence national contributions to warming to a similar degree as the inclusion of N₂O. Note that N₂O-related warming contributes around 7% of the warming related to all three GHGs in this analysis on average across individual countries (standard deviation 5%). In addition, the cooling effect of sulphate aerosols and other reflective aerosol species is not included here, yet we note that the cooling effect of aerosols is comparable in magnitude to the warming effect of CH₄ at the global scale^{1,8}. Consequently, changes in national contributions to warming would occur if other gases and aerosols were to be included in this analysis. For example, including aerosols has been estimated to reduce China's contribution to warming to 8%, as compared with 11% in a case including only well-mixed GHGs²⁴.

Second, contributions to warming depend on the time period under consideration. For example, the inclusion or exclusion of pre-industrial LULUCF CO₂ emissions has a small influence on the contribution made by countries whose key period of land use change preceded the industrial period (up to a few percentage points in European countries and China)¹⁶. Here, we consult multi-gas emissions datasets that collectively include the years 1851–2021, and we report on contributions to climate change since 1850 (note that the CH₄ emissions data for years 1830–1849 are also required to calculate cumulative CO₂-equivalent emissions from 1850 onwards, see **Methods**). Figures 5, 6 show how national contributions to warming have evolved with time since 1850. However, we note that earlier or later reference years would provide a different perspective on national contributions to emissions. For example, selecting a reference year of 1900 would reduce cumulative global CO₂ emissions by 40 Pg CO₂ and lessen the related warming by 0.02 °C (–2.3%). For national contributions, the corresponding effect of varying the reference year on warming would depend on the fraction of cumulative national emissions that occurred before or after the reference year for any particular country. A change in reference year within 1850–1900 has a considerably smaller impact on the GMST responses to global or national CH₄ emissions due to lesser dependence of CH₄-related warming on cumulative emissions than in the case of CO₂ or N₂O.

Third, contributions to warming depend on population. We do not include per capita emissions or per capita contributions to warming in our dataset. Nonetheless, we note that previous work has highlighted per capita expressions of emissions or warming as a means of accounting for differences in the intensity of emissions or warming impact per country, providing further perspective on the national accountability for climate change^{3,12}.

Finally, contributions to warming depend on international trade. Some countries (e.g. China and India, and Brazil) emit CO₂ in the process of producing goods or services for export (in *net* terms), while other countries/regions (e.g. the EU27 and the USA) are *net* importers and consume goods or services which require emissions in external territories. Here, we do not account for national emissions embodied in goods or services traded between countries (i.e. the emissions estimates used here include territorial emissions only rather than consumption-based emissions). Estimates of consumption-based emissions are available for fossil CO₂³³ and for LULUCF CO₂, CH₄ and N₂O⁹⁸ and could be used to produce consumption-based national warming contributions, however these records begin only in the 1960s–1970s.

Code availability

The R Statistics code used to perform all methods described here can be accessed via the GitHub repo at the following link: https://github.com/jonesmattw/National_Warming_Contributions.git.

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Competing interests

The authors declare no conflict of interest.

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